



Using W phase for regional tsunami warning and rapid earthquake hazard assessment

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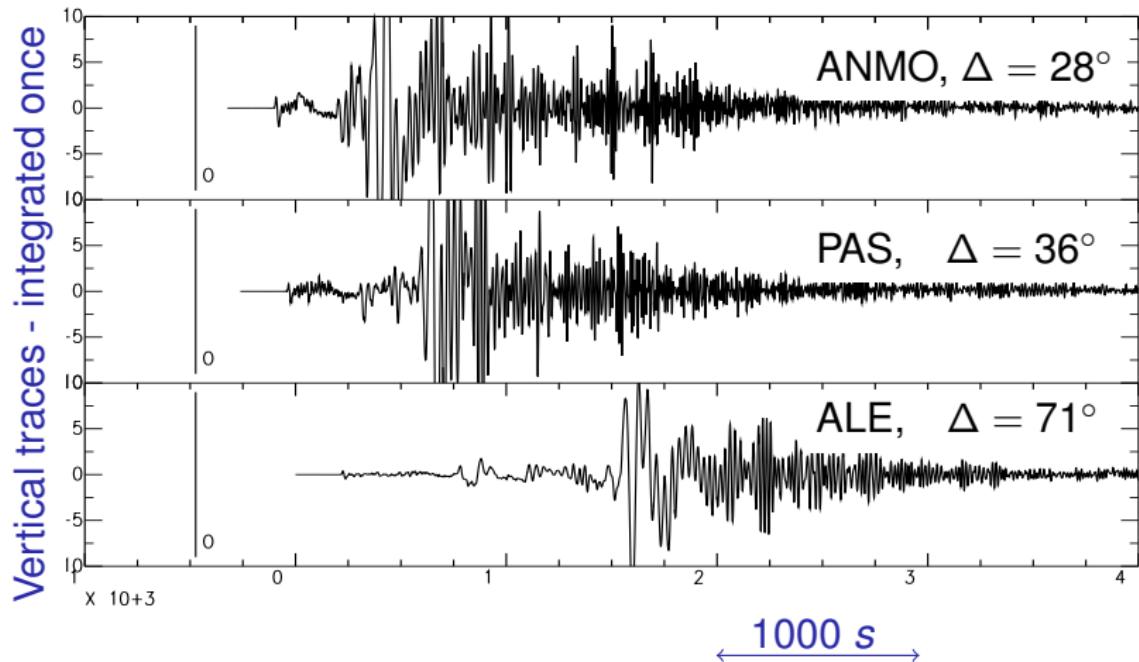


Introduction

- ▶ We have recently developed a source inversion technique based on the waveform modeling of W-phase.
- ▶ W-phase is a very long period (200s-1000s) phase arriving right after the P.
- ▶ It was first recognized after the 1992 Nicaragua earthquake.
- ▶ The inversion technique was originally devised to work for large events ($M_w > \sim 7.5$) with teleseismic data and it provides a VLP characterization of the source (e.g. Tsunami earthquakes).
- ▶ **We explore here the possibility of an application with regional data and with smaller magnitudes.**

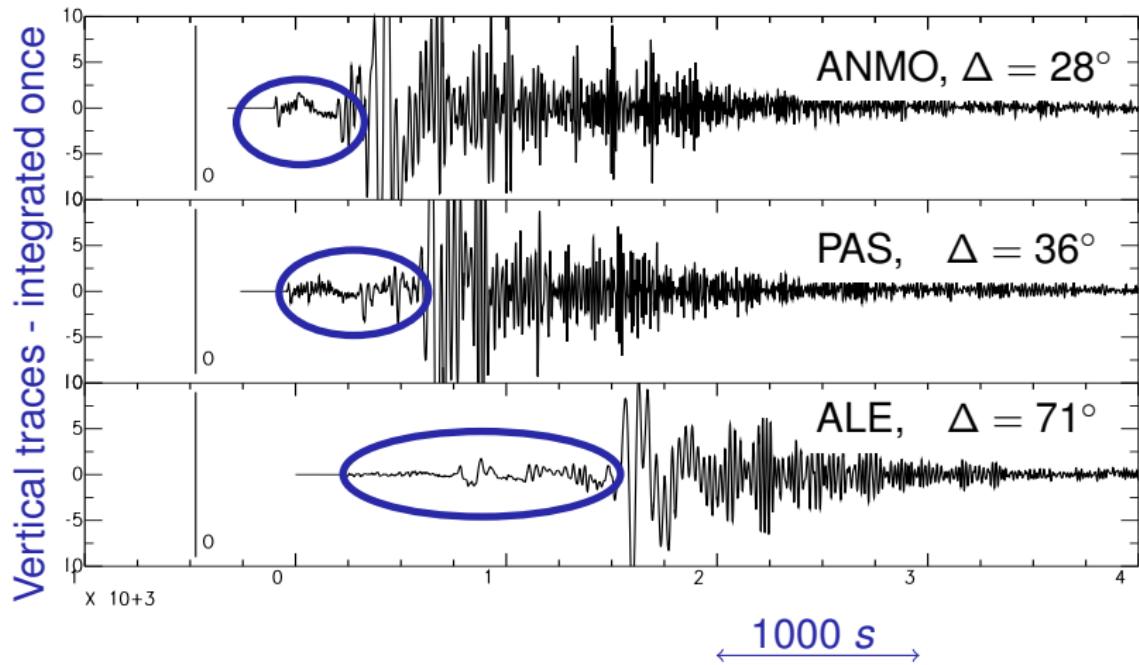
W-phase

Example: Nicaragua, 1992



W-phase

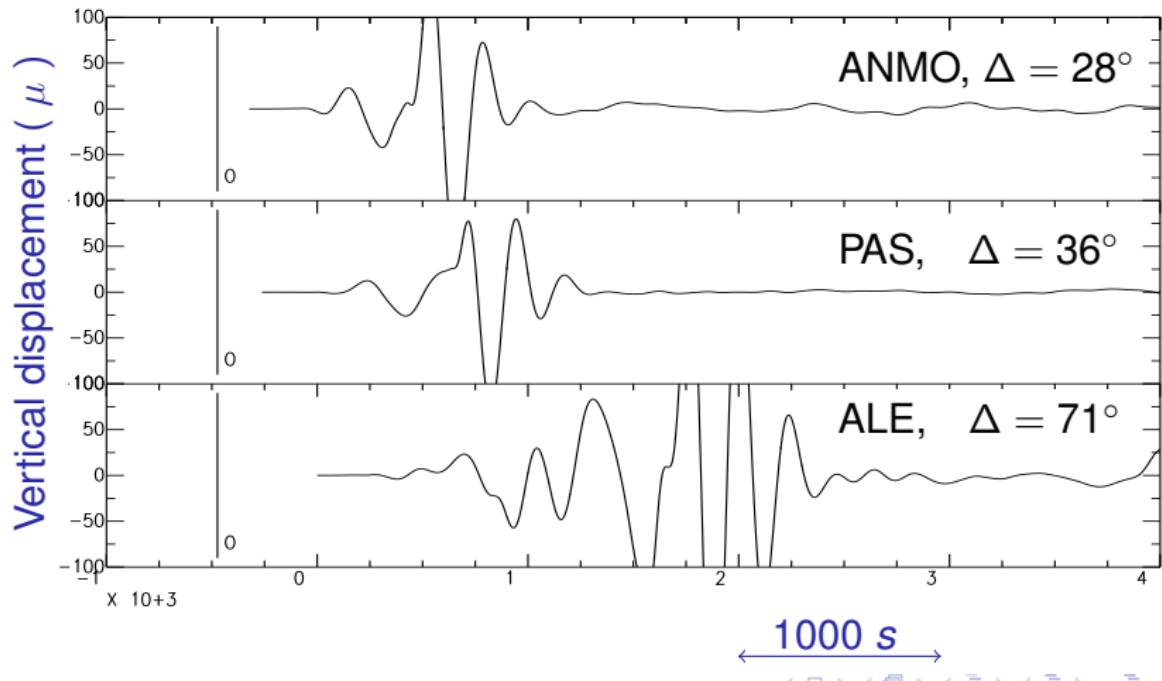
Example: Nicaragua, 1992



W-phase:

Example: Nicaragua, 1992

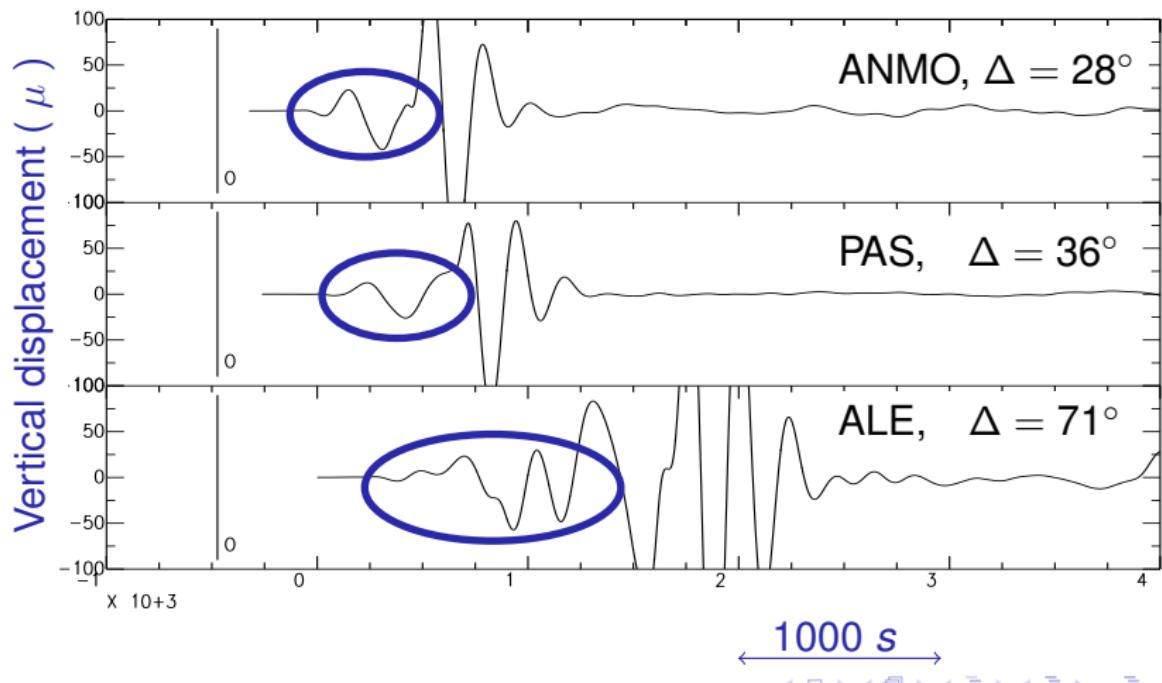
Deconvolved + bp 200s-1000s



W-phase:

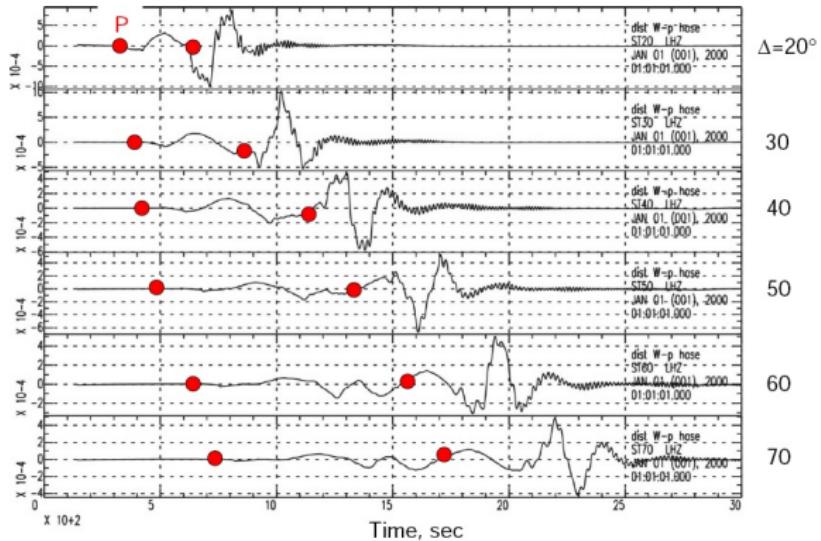
Example: Nicaragua, 1992

Deconvolved + bp 200s-1000s

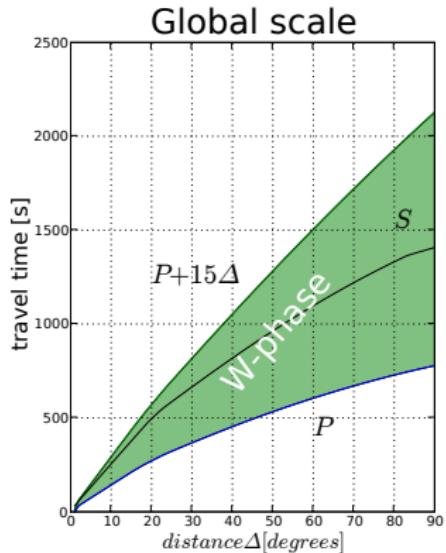


Time window

W-phase time-window
 $P, P + 15\Delta s/o$



W-phase in the global context



Definition

- ▶ Time window: $P, P+15\Delta$
- ▶ Bandpass: .001 Hz - .005 Hz

Properties

- ▶ Fast group velocity: 4.5-9 km/s
- ▶ Fairly insensitive to:
 - ▶ shallow lateral heterogeneities
 - ▶ source second order details
- ▶ Avoid large amplitude surface waves

Source retrieval from W-phase

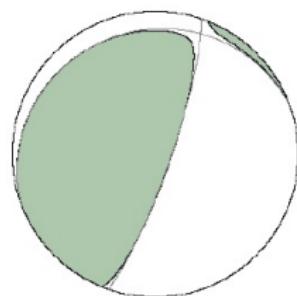
Inversion: main features

- ▶ Time domain
- ▶ Point source (VLP data)
- ▶ Need a preliminary source location: PDE, JMA
- ▶ Library of precomputed Green's functions
- ▶ Linear inversion → Moment tensor components
 - ▶ PDE ($\Delta < 50^\circ$): $t_0 + 20$ min
 - ▶ Grid search ($\Delta < 90^\circ$): $t_0 + 35$ min
- ▶ RT implementation: β -test at NEIC-USGS, (Gavin Hayes)

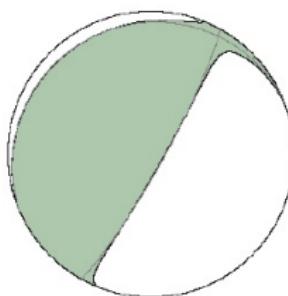
Example: global data

Tokachi-Oki-2003

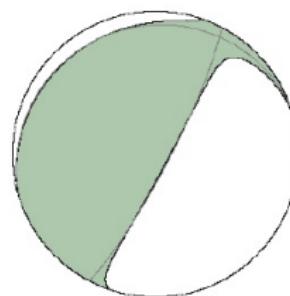
2003 Tokachi-oki WP inversion



PDE location
 $t_h = 30s, t_d = 30s$
 $M_w = 8.24$



Optimized centroid
 $t_h = 30s, t_d = 30s$
 $M_w = 8.31$



GCMT centroid
 $t_h = 31.8s, t_d = 33.5s$
 $M_w = 8.27$



CMT
 $M_w = 8.3$

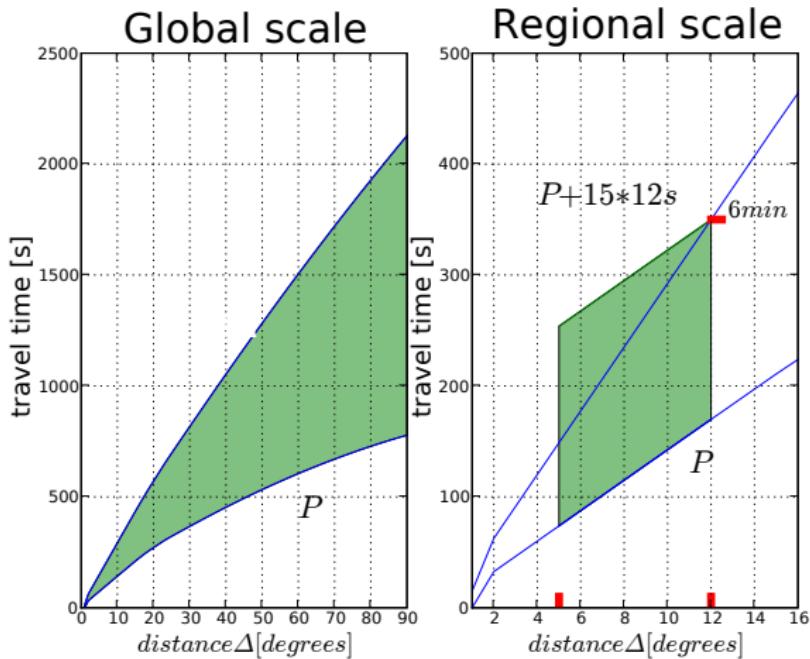
$t_0 + 20 \text{ min}$ $t_0 + 35 \text{ min}$

Regional data

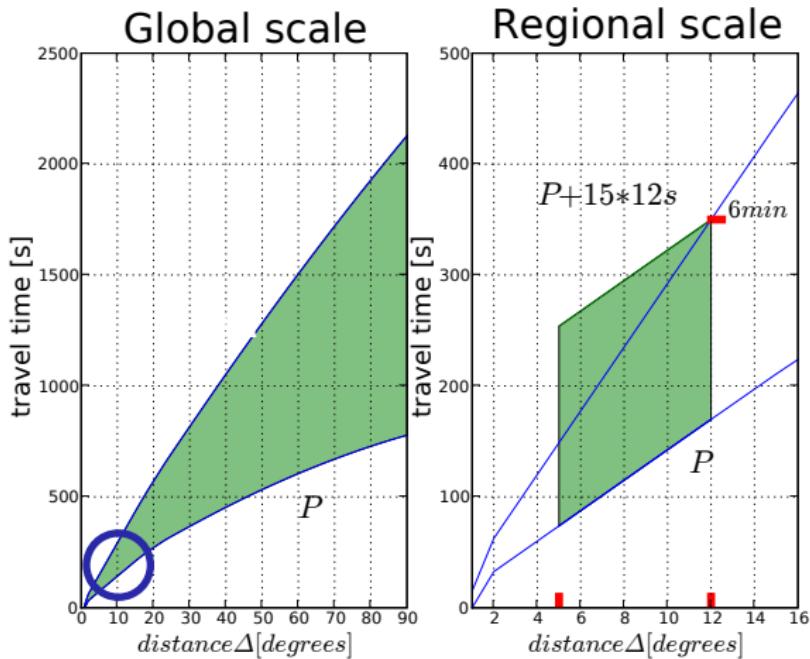
Extension to regional data and lower magnitudes

- ▶ Target: $M_w > \sim 6.5$
- ▶ Data distribution:
 - ▶ $\Delta < 12^\circ \rightarrow 6 \text{ min.}$
 - ▶ $\Delta > 5^\circ$: high gain data, nonlinearity.
- ▶ Modifications:
 - ▶ Time window: $(P, P+15\Delta)$ inappropriate
 - ▶ Frequency band: signal/noise ratio

W-phase time window

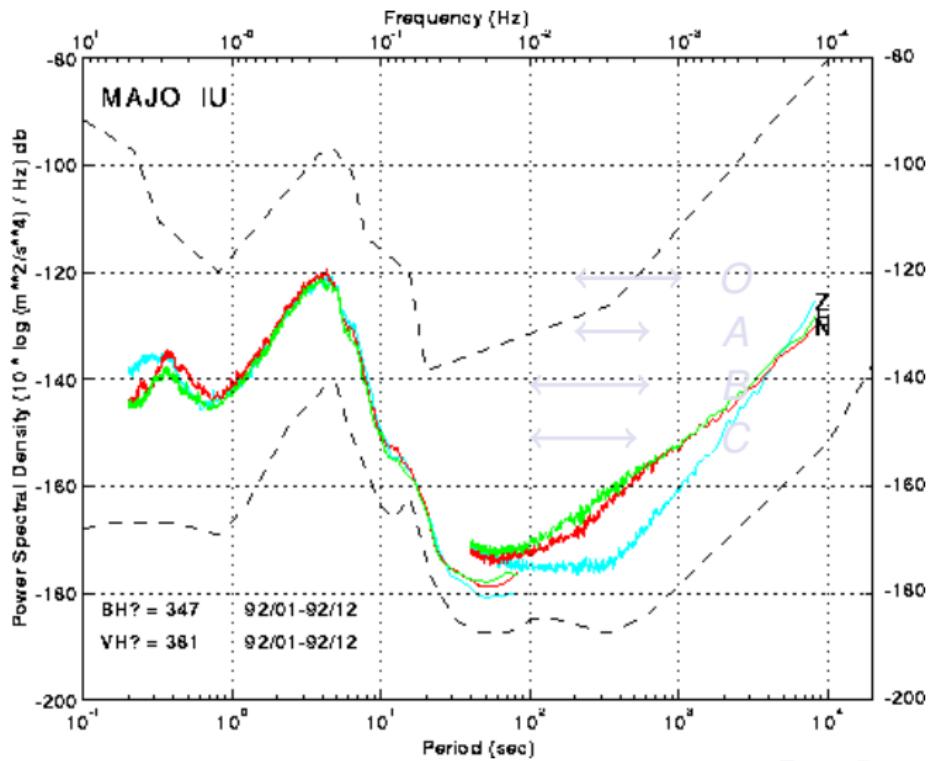


W-phase time window



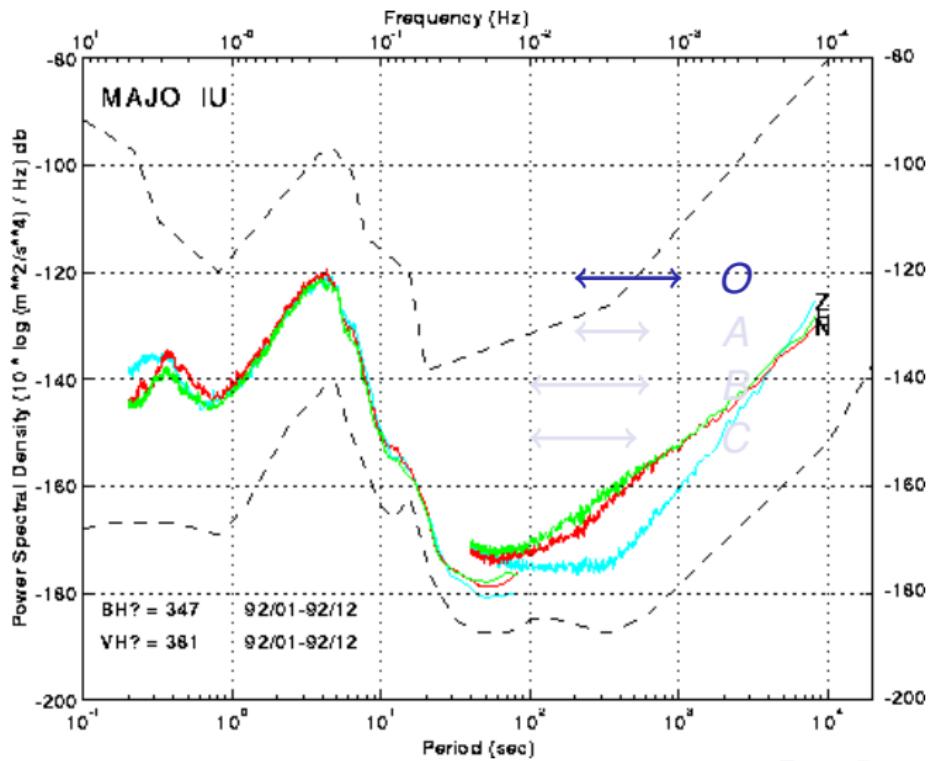
Extension to lower magnitudes

Acceleration noise spectrum at MAJO



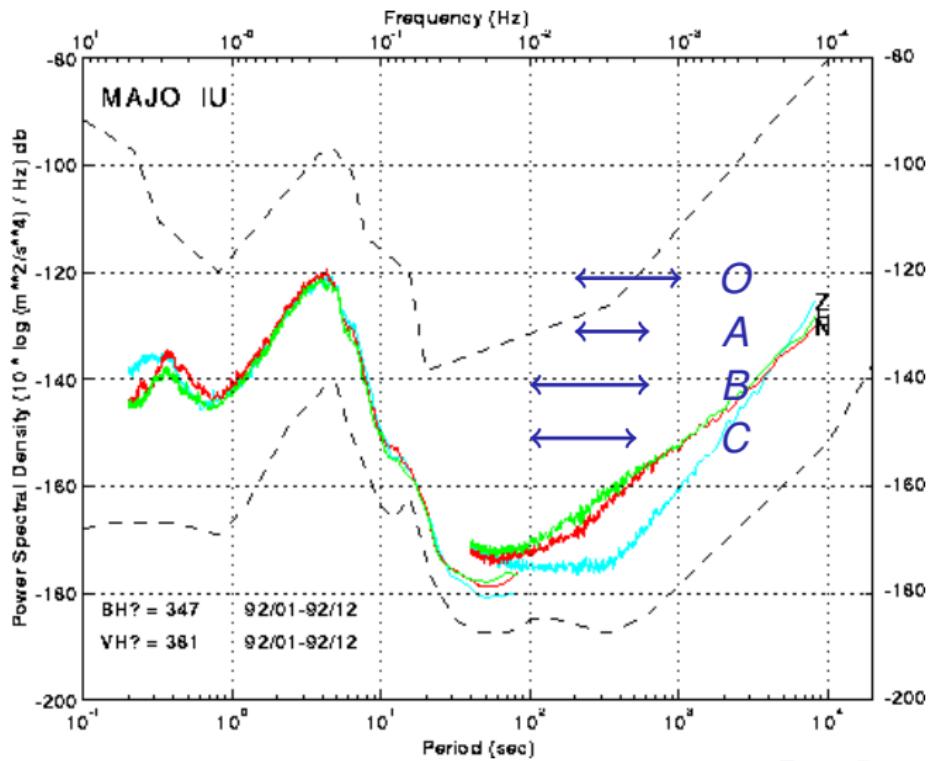
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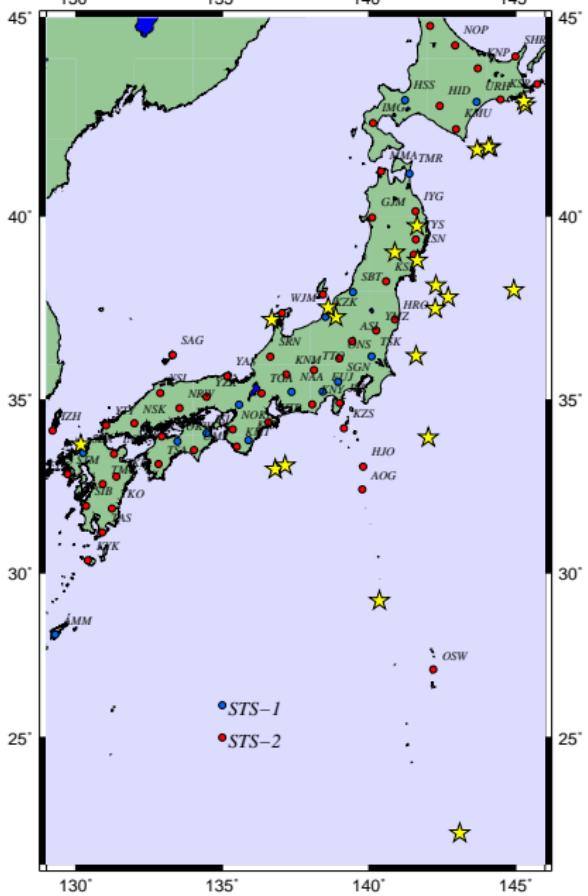
Data:

Japanese broadband network (F-net)

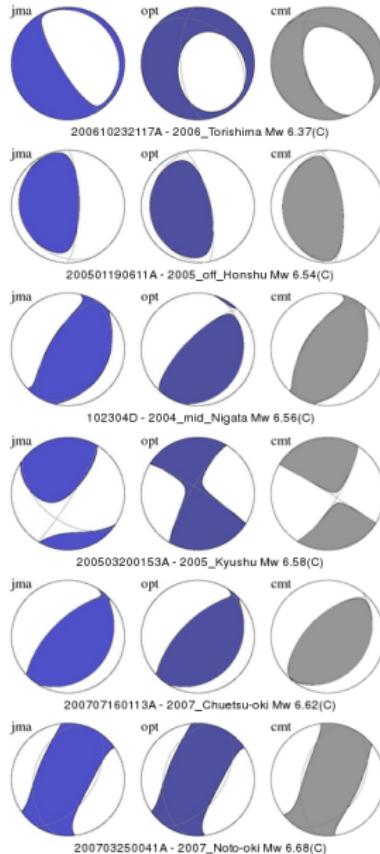
Events:

$$M_{jma} > 6.7$$

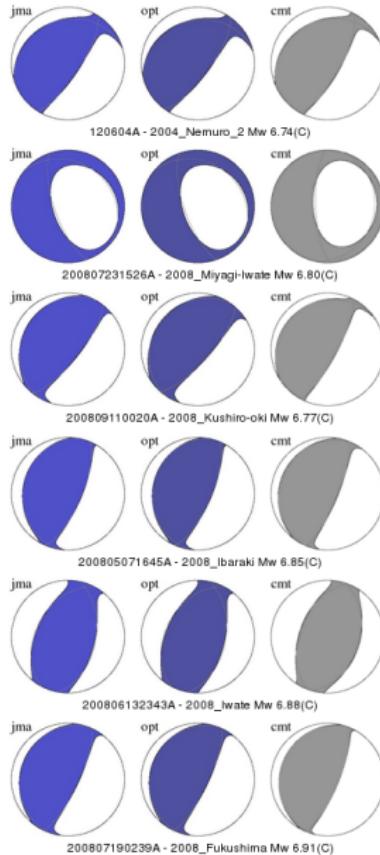
2003-2008



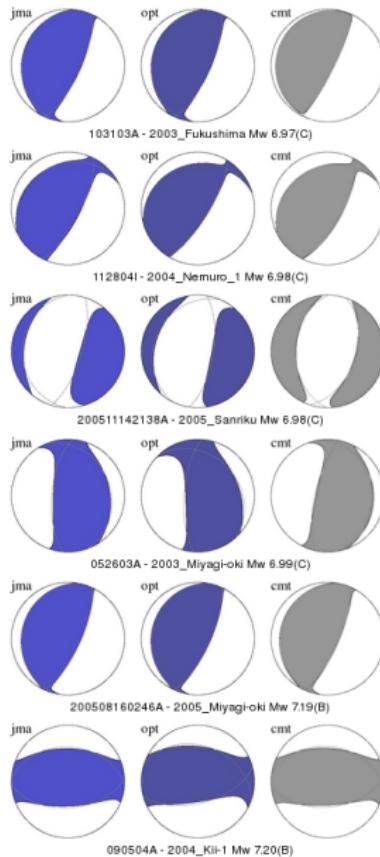
F-net, 2003-2008, $M_{jma} > 6.7$ (1/4)



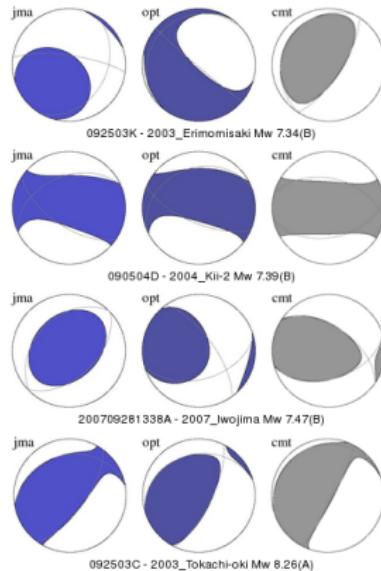
F-net, 2003-2008, $M_{jma} > 6.7$ (2/4)



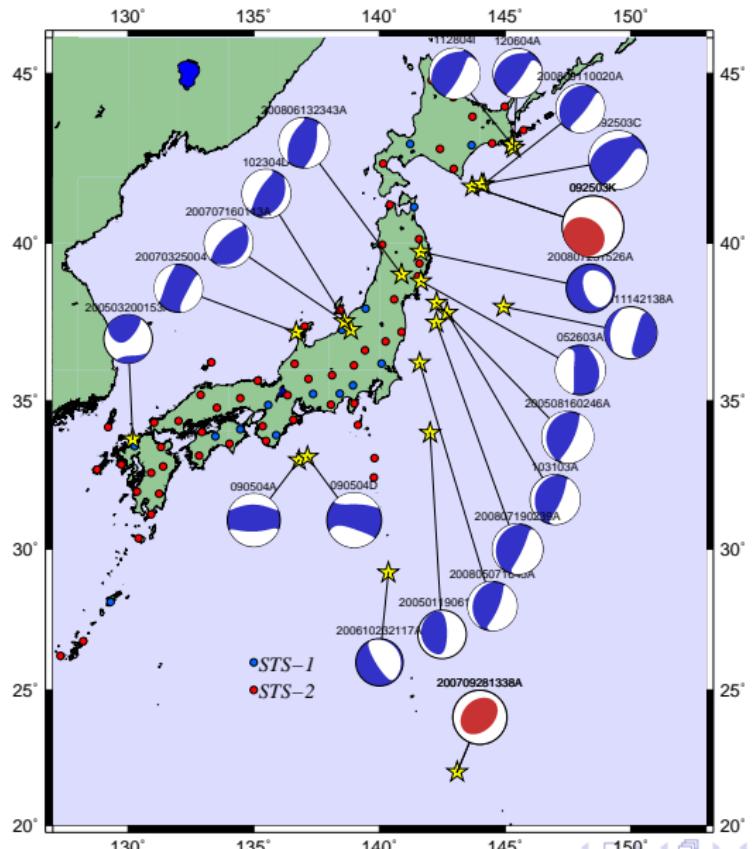
F-net, 2003-2008, $M_{jma} > 6.7$ (3/4)



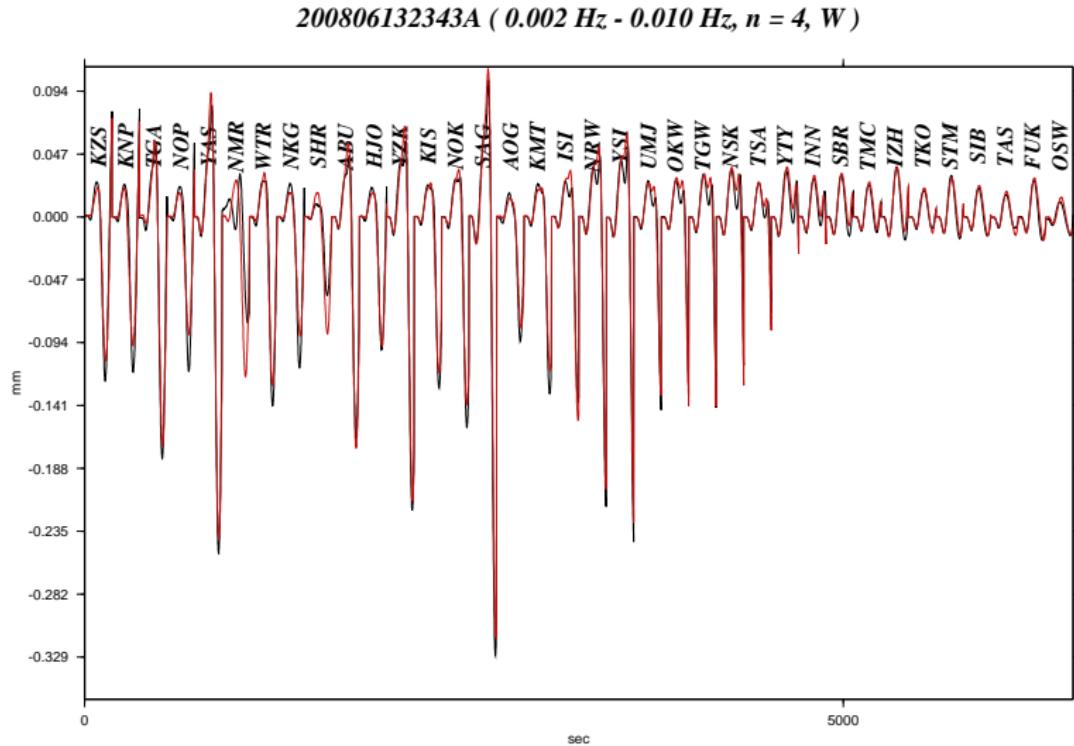
F-net, 2003-2008, $M_{jma} > 6.7$ (4/4)



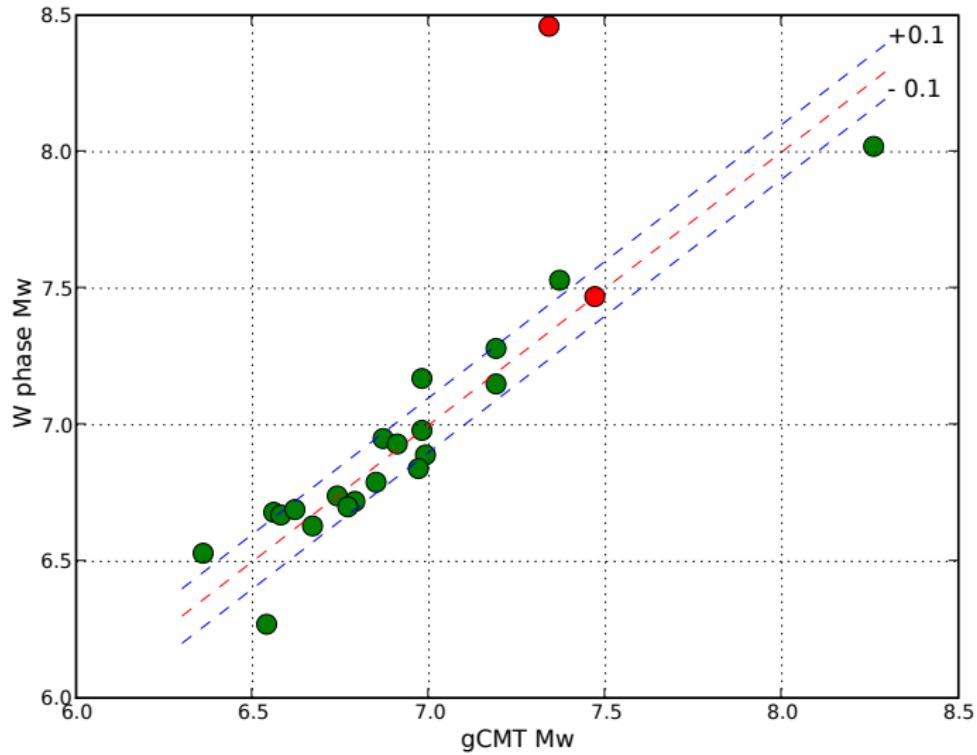
Regional W phase focal mechanisms ($t_0 + 6\text{min}$)



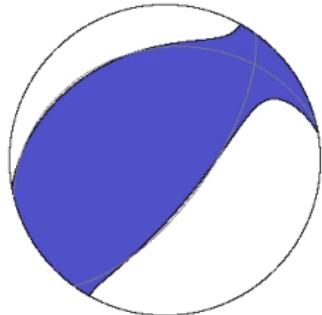
Regional W-phase, example of fit: 2008 Iwate



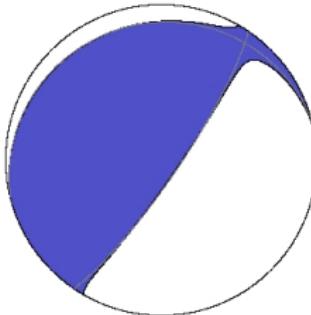
Moment Magnitude: gCMT - W-phase



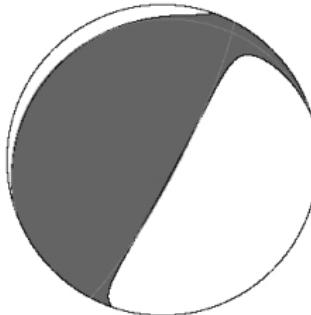
2003 Tokachi-oki: depth effect



$H = 45\text{km}$
 $\delta = 29^\circ$
 $M_w = 8.02$



$H = 28\text{km}$
 $\delta = 14^\circ$
 $M_w = 8.15$



$HVD(H = 28\text{km})$
 $\delta = 11^\circ$
 $M_w = 8.26$

Conclusions

- ▶ We use F-net data in the range ($5^\circ < \Delta < 12^\circ$)
- ▶ Time window: $t_P, t_P + 180\text{s}$
- ▶ Variable frequency band:
 $(.00167\text{Hz} - .005\text{Hz}) \rightarrow (.005\text{Hz}, 010\text{Hz})$
- ▶ Moment tensor solution available at $t_0 + 6\text{min}$
- ▶ Can be done completely automatic and
- ▶ Provide a solution 6 min after the origin time.